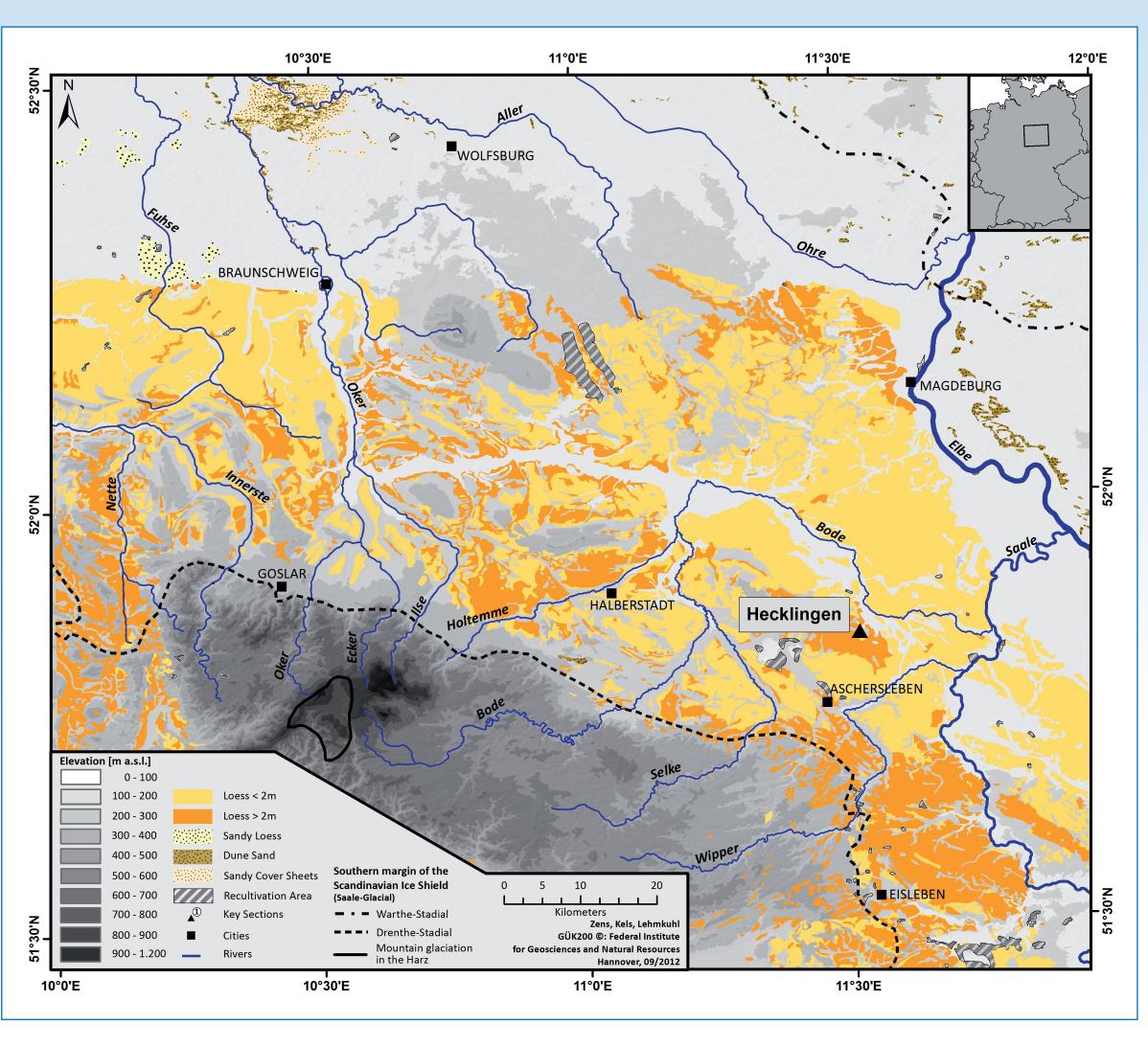
Environmental Conditions during the Last Glacial Cycle in the Northern Harz Foreland

Lydia Krauß, Jörg Zens, Christian Zeeden & Frank Lehmkuhl

1. Introduction

For the reconstruction of past environmental and climatic conditions several archives can be investigated, e.g. for the global scale deep-sea sediment cores¹ or polar ice cores² and for a regional scale lake sediment cores³. One important type of regional terrestrial archives of the Quaternary are loesspaleosol-sequences (LPS)⁴.

Loess is an aeolian, mainly silt sized (< 63 µm), usually calcareous and yellowish loose sediment which occupies 10 % of the land surface of the Earth. During cold and dry phases loess accumulates. On it soils develop during warmer periods. Through cyclic changes of climatic conditions a series of loess



Within the CRC 806 – Our Way to Europe – that studies the dispersal of the modern Homo sapiens from Africa to Europe, we analyze LPS in Central Europe to reconstruct environmental conditions during the last glacial cycle, to find out if the environmental conditions were too rough for humans to settle here⁹.

The Northern Harz Foreland represents a region that was close to the ice sheet during the last glacial cycle¹⁰. Here environmental conditions were probably pretty harsh, but to find that out, we analyzed a LPS in Hecklingen (fig.1).

and paleosols develop on each other^{5,6}.

In comparison to lake sediments and ice cores LPS are archives close to archeological findings. Early humans walked, settled and lived on the soils developed on loess as we do today. This is no wonder because those soils have optimal conditions for vegetation to grow abundantly⁷.

However, the modern humans arrived in Central Europe around 33,000 years ago for the first time, but disappeared again till around 14,000 years ago. After that they were permanently present in Central Europe⁸.

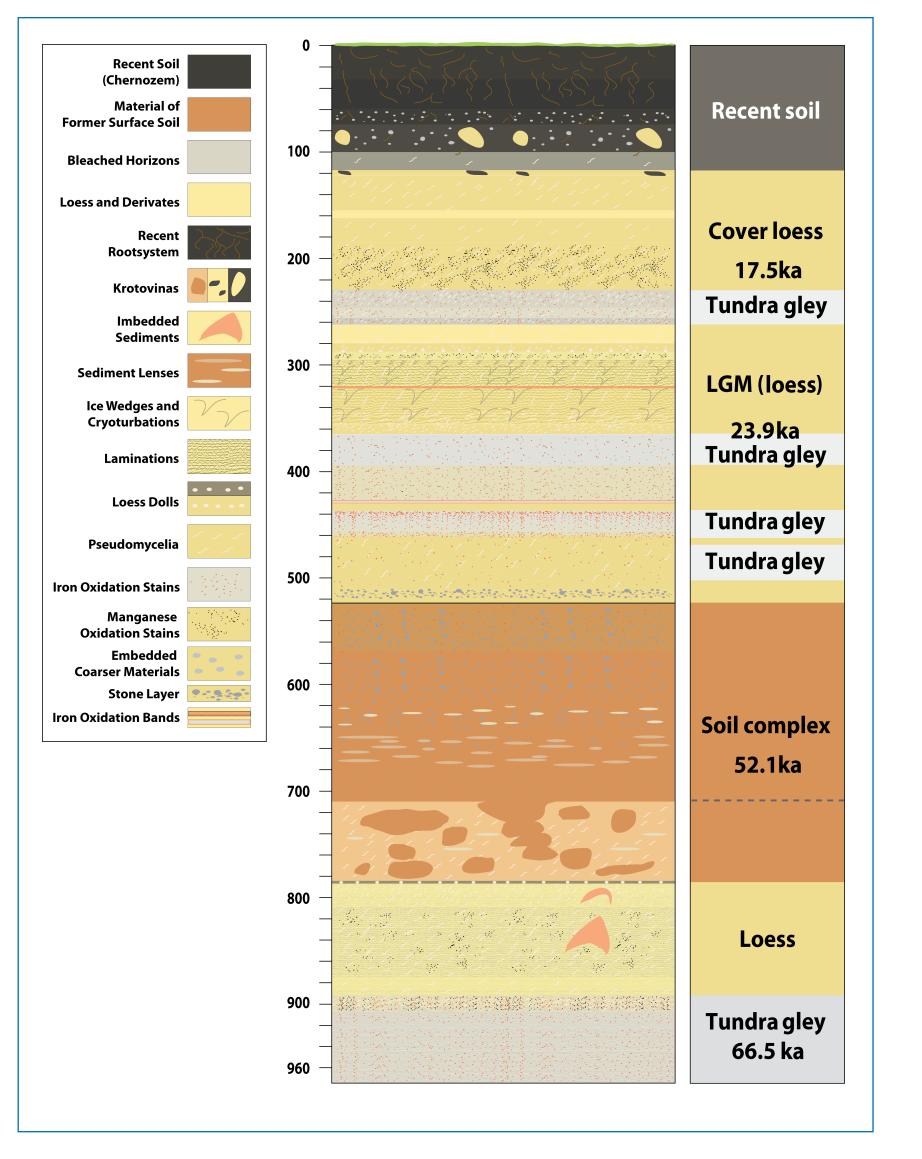


Figure 1: Map of the Northern Harz Foreland with the loess distribution

2. Features of our profile

In Hecklingen we identified a recent soil of 1.2 m thickness, three around 30 cm thick tundra gleys with a typical grey matrix and red oxidation stains in 2.3, 3.6 and 4.3 m depth, an almost 2 m thick red-brownish soil with an additional 70 cm thick transition layer with krotovinas and an at least 80 cm thick tundra gley at the bottom of the sequence (fig.2).

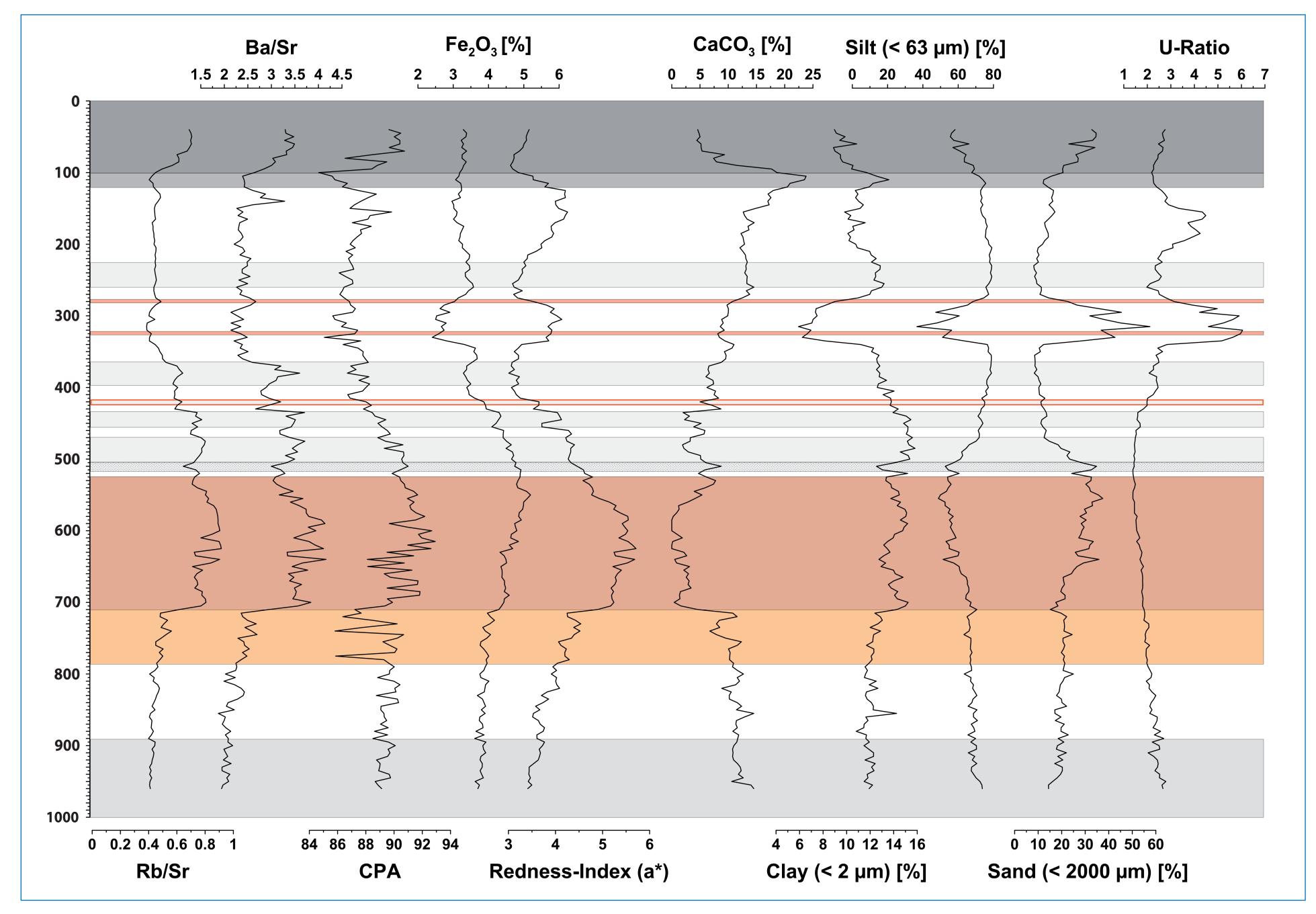


Figure 2: Detailed and generalized profile drawing with OSL ages after Reinecke (2006)¹⁰

In between those layers loess was embedded. The lowermost loess layer below the red-brownish soil material showed lamination. The 1 m thick loess layer below the first tundra gley was laminated, too, but additionally showed strong cryoturbations and ice wedges in some areas. Within the uppermost loess layer no lamination but pseudomycelia and manganese oxidation stains were visible (fig.2).

Additionally, we look at the grain size distribution within the profile, e.g. if we have a lot of clay present, we can determine soil development, since most clay minerals are produced by pedogenic processes. If there is an increase in silt content, we can assume an aeolian material input with a long distance material transport. This speaks for dryer and colder conditions. Higher sand contents on the

The grain size distribution show three phases of higher sand content, speaking for a closer source of the here accumulated material. The U-Ratio show two phases of increased aeolian material input (fig.3).

Acknowledgements

This study was financially supported by the German Research Foundation (DFG) in the frame of the Collaborative Research Center Our way to Europe (CRC 806).

Figure 3: Results of geochemical, color and grain size measurements

3. What do we analyze and why

For the reconstruction we use different proxies. The geochemical and

color parameters show how strong the material is weathered. Is the sediment strongly weathered, we can assume that we have soil

development. Further, this is indicating warmer and more humid climatic conditions.

other hand indicate a close-by source input.

4. A brief summary of the results

The profile shows strong disturbances, here especially lamination

and cryoturbation, which are common in central Europe and make

the interpretation of the profile more challenging (fig.2).

The weathering indices, e.g. Rb/Sr, Ba/Sr and CPA, show higher values in the identified soils and soil like material, supporting the field observations (fig.3). References

¹L. E. Lisiecki & M. E. Raymo (2005): A Pliocene-Pleistocene stack of 57 globally distributed benthic δ¹⁸O records. Paleooceanography, Vol. 20, PA1003.

²S. J. Johnson, D. Dahl-Jensen, N. Gundestrup, J. P. Steffensen, H. B. Clausen, H. Miller, V. Masson-Delmotte, A. E. Sveinbjörnsdottir & J. White (2001): Oxygen isotope and palaeotemperature records from six Greenland ice-core stations: Camp Century, Dye-3, GRIP, GISP2, Renland and NorthGRIP Journal of Quaternary Science, Vol. 16 (4), pp. 299–307.

³F. Sirocko, K. Seelos, K. Schaber, B. Rein, F. Dreher, M. Diehl, R. Lehne, K. Jäger, M. Krbetschek & D. Degering (2005): A late Eemian aridity pulse in central Europe during the last glacial inception. Nature, Vol. 436, pp. 833-836.

⁴M. Frechen, E. A. Oches & K. E. Kohfeld (2003): Loess in Europe—mass accumulation rates during the Last Glacial Period. Quaternary Science Reviews, Vol. 22(18–19), pp. 1835-1857.

⁵K. Pye (1996): The nature, origin and accumulation of loess. Quaternary Science Reviews, Vol. 14, pp. 653–667.

⁶S. R. Taylor, S. M. McLennan & M. T. McCulloch (1983): Geochemistry of loess, continental crustal composition and crustal model ages. Geochimica et Cosmochimica Acta, Vol. 47(11), pp. 1897-1905.

⁷D. Haase, J. Fink, G. Haase, R. Ruske, M. Pécsi, H. Richter, M. Altermann & K.-D. Jäger (2007): Loess in Europe — its spatial distribution based on a European Loess Map, scale 1:2,500,000. Quaternary Science Reviews, Vol. 26 (9-10), pp. 1301-1312.

⁸http://www.sfb806.uni-koeln.de/index.php/projects/cluster-d/d1

9 http://www.sfb806.uni-koeln.de/

¹⁰V. Reinecke (2006): Untersuchungen zur mittel- und jungpleistozänen Reliefentwicklung und Morphodynamik im nördlichen Harzvorland. Aachen (= Aachener Geographische Arbeiten 43).

Lydia Krauß Physical Geography and Geoecology Department of Geography, RWTH Aachen University +49 241 80 96052



